

Solar System Science with the TMT

Glenn Orton Jet Propulsion Laboratory California Institute of Technology

Thanks To Leigh Fletcher, as well as Athena Coustenis, Therese Encrenaz, with additional input from N. Bowles, T. Greathouse, R. Giles, C. Howett, P. Irwin, R. de Kok, H. Melin, I. de Pater, T. Stallard, C. Tsang, N. Teanby, C. Wilson



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TMT: New Opportunity for Planetary Science

- Spatially-resolved imaging and spectroscopy is a crucial diagnostic of environmental conditions within our solar system.
- Understand \bullet diversity of planetary climates and processes to place our own world in broader context.
- Combination of exquisite spectral/spatial resolution enables new opportunities for planetary science.



Intro: Spectral range?

- TMT instrumentation should cover both reflected sunlight ($<5 \mu m$) and thermal emission ($>5 \mu m$) in the IR.
 - Relate climate/environmental conditions to surface/atmosphere albedo characteristics.
- For atmospheres, IR provides:
 - **third dimension** (vertical) to otherwise 2D images.
 - IR signatures of traces gases used to understand planetary chemistry and circulation.
- IR allows study of energy exchange processes (aurora, ionospheres).
- IR provides access to elemental and isotopic composition of primitive material to test **planetary origins**.
- Thermal infrared has been overlooked over the past decades in favor of reflectance spectroscopy.
 - Thermal-IR (> 5 μ m) absent from future outer solar system missions.
 - Infrared spatial resolution and spectral coverage primitive compared to 'optical/near-IR' wavelengths.
 - TMT will help change that!





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Intro: High Spatial Resolution

- TMT diffraction-limited seeing factor of 6 improvement over JWST, provided good AO correction.
- Issues for largest targets:
 - Small FOV need mosaics?
 - Chopping off disc?
 - AO guide star availability and angular separation?
 - Brightness and saturation?



Io in L-band, Keck/NIRC2, Mignant and Marchis, 2002.

Size of Planetary Targets

Target	Target Diameter		Diameter
Venus	66"	Neptune	2.4"
Mars	25.1"	KBO/Pluto	0.11"
Ceres	0.84″	lo	1.2″
Vesta	0.64"	Ganymede	1.8″
Jupiter	50.1"	Titan	0.8″
Saturn	20.1"	Triton	0.13″
Uranus	4.1″		

Diffraction-Limited Resolution

Telescope	JWST	VLT	TMT
Wavelengh (µm)	6.5	8	39
3	0.116	0.094	0.019
4	0.155	0.126	0.026
5	0.194	0.157	0.032
8	0.310	0.252	0.052
14	0.542	0.440	0.090



Giant Planet I: Climate & Circulation

Variation of climates

- What is the 3d circulation of giant planet atmospheres and the factors shaping variability?
- Aim to **connect visible changes** (clouds, colors, winds) to **environmental changes** (temperatures, composition, wind shear).
- 3D Reconstruction of atmospheric temperatures, composition, winds, clouds, etc.
- Suited to imaging and moderate-resolution spectroscopy, in the L- through Qbands.

How does atmospheric circulation vary between planets?



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Giant Planet II: Temporal Changes

• Giant planet atmospheres evolve on timescales from:

1989 AUG 16 8 UT

- Minutes (impacts)
- Days (Plumes, storms)
- Weeks (Belt/zone life cycles)
- Years (seasonal evolution)
- The TMT would provide the same understanding of ice giants as we have for Jupiter/Saturn.
- Environmental changes underlying this variability.
 - Would benefit from simultaneous measurement of reflected sunlight and thermal imaging
- TMT diffraction-limited resolution provides good altitude coverage for temperatures, gaseous abundances and clouds
 - Temperatures from N- and Q-band H₂ absorption and CH₄ emission imaging and spectroscopy
 - Gaseous abundances from N-band thermal emission, primarily from hydrocarbons
 - Cloud properties from <5 µm reflection spectroscopy, K-, M-, N- and Q-band center-to-limb behavior

7.8 µm Keck LWS

What drives atmospheric variability?

Variation of

climates

Saturn's thermal emission Orton & Yanamandra-Fisher 2005

VLT / VISIR 2006 Neptune's thermal emission, Orton et al., 2007



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Giant Planet III: Deep into the Clouds

Fundamental Processes

- M-band provides a unique window minimal reflected sunlight, minimum H_2 and CH_4 opacities allow deep sounding.
- Giant-planet cloud condensation region.
- M-band spectroscopy allows separation of gaseous species.
 - NH₃ for cloud-condensation
 - PH₃, AsH₃, GeH₄ for deep atmospheric circulation
 - CO for external influx of O-species.
 - CH_3D for cloud sounding, D/H ratio for origins.
 - High-temperature H_3^+ emission
- Never exploited for the ice giants before!







Vortices at 5 µm from Keck. Credit: Imke de Pater, Michael Wong (UC Berkeley); Al Conrad (Keck), and Chris Go (Cebu, Philippines)

NASA

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Giant Planet IV: Planetary Aurora

Energy exchanges

How does solar wind control planetary aurorae and heating?

- H_3^+ auroral emission allows mapping of giant planet ionospheres (temperature/density) by tracing heat flow within upper atmospheres.
 - Ionospheric flows coupled to magnetic field, connects planet to external environment.
 - Focus on auroral variability; but....
 - Non-auroral regions are driven by complex dynamics that comes from multiple sources, including low-latitude precipitation, upper atmosphere dynamics and upwelling heat
 - How does upper atmospheric temperature change with altitude/latitude – a key question for planetary aeronomy!
- Cassini can image Saturn's aurorae, but low spectral resolution prevents measurements of ionospheric properties
- Uranus H₃⁺ hard to detect (first seen in 1992), but long-term upper atmospheric temperatures have decreased over time.
- Neptune H₃⁺ still undetected, models predict it should be there!
- L-band measurements are required.



Saturn's ring-rain, O'Donoghue et al., 2013

H₃⁺images of Jupiter's aurora





NASA

L - band (3.5 - 4.0 µm)

06:18 UT



•

Giant Planet V: Planet-Satellite Interaction

How are planetary ionospheres coupled to their satellite systems?

- Plasma sources within the magnetosphere (e.g. Io) cause auroral UV footprints, but also IR heating.
 - How far does heating extend, vertically & latitudinally?
 - How is the energy dissipated with time?
 - Does Enceladus produce a similar spot at Saturn?
 - Do auroral spots vary with volcanic (Io) or water plume (Europa, Enceladus) activity?
- Insights into energy flow between ionosphere, magnetosphere and satellites.
- Requires the high (i) sensitivity, (ii) spectral resolution and (iii) spatial resolution.

Hubble UV imaging of auroral footprints.



Saturn's IR aurora from Cassini/VIMS



Giant Planet VI: Middle Atmosphere

Fundamental Processes

- Wave activity may dominate energy transport and material distribution in planetary stratospheres.
- Investigate via:
 - (i) High-resolution sounding of Doppler broadened CH₄, C₂H₆ lines (3, 7, 12 μm)
 (ii) Stellar occultations to probe atmospheric T/density structure (e.g., using Io hotspots as point source)



Saturn's SAO from Fouchet et al., 2008



Jupiter occultation study, Greathouse et al., 2010



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Terrestrial I: Trace Species (Venus)

Variation of climates



Variability of SO₂ mapped using ratios of SO₂/CO₂ lines at 7 μ m, Encrenaz et al., 2013

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Terrestrial II: Trace Species (Mars)

What are the sources and sinks of trace species in the Martian atmosphere?

- Sensitive search for trace species on Mars (e.g., controversial Martian methane).
 - CH₄ at 3.3, 7.8 μm
 - HDO at 3.7 μ m, unsuccessful searches by ISO.
- Non-LTE emission from CO and CO_2 first detected by ISO (Lellouch et al., 2000).
 - Probing upper atmospheric circulation, as on Venus.
- Carbonate signatures in the M and N bands tentatively identified by Lellouch et al. (2000).







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Fundamental

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Variation of climates

Titan's rich atmospheric \circ composition considered as an example of a prebiotic atmosphere.

assini

- Extensive studies by Cassini/VIMS and Cassini/CIRS will come to an end in 2017.
 - TMT could continue these over a full Saturnian year.
- High-resolution imaging (reflectivity and thermal) and N-band spectra to understand:
 - Methane cycle and variable cloud activity (K- and L-band)
 - Seasonal evolution of circulation through temperatures, trace gases and cloud activity (N-band)

2 Sep. 2004 2 Oct. 2004 3 Oct. 2004 7 Oct. 2004 3 Nov. 2004 Variable cloud activity monitored by Keck in reflected sunlight, Roe et al., 2008. Tracer 500 km enriched Meridional circulation descending branch air VORTEX Tracer ~ 1 zonal wind contour depleted mesospheric Mixing of long zone 300 km lived species Altitude Residual Upward advection circulation Barotropic instabilities 100 km Condensation Titan's atmospheric circulation during northern winter, Teanby Meridional circulation returning branch 50 km -50° 'n 65° -80

Latitude

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et al., 2008

'Terrestrial' IV: Titan Surface to Upper Atmosphere

Fundamental Processes

- Surface variability from M-band: Using high-resolution infrared windows (e.g., M-band where CH₄ opacity at a minimum).
- Upper atmospheric studies from:
 - CO fluorescence near 4.7 µm allow measurement of external oxygen influx.
 - HCN fluorescence near 3 µm to understand upper atmospheric circulation.





Surfaces I: Volcanic Io

Surface environments

- Unprecedented resolution will allow us to monitor volcanism on a distant world.
- Io thermal output finding the missing power
 - Currently miss smaller volcanoes/lava flows, missing maybe halfheat output.
 - Cooler flows will show up at longer N-band and Q-band wavelengths.
- Surface atmosphere interaction SO₂ frost:
 - SO₂ atmospheric collapse in eclipse should strengthen surface ice bands 1.9-4.2 μm, but never seen, K-, L- & M-band spectroscopy.



Marchis et al., 2005 using Keck L and M band imaging





Surface environments

Surfaces II: Europa

Can icy surfaces reveal insights into their watery interiors and plasma interactions?

- Reading surface geology and composition reveals (i) connection to potentially-habitable subsurface ocean; and (ii) effects of interaction of surface materials with plasma environment.
- Search for evidence of endogenic emission associated with plume/geologic activity (as on Enceladus).
 - Location of plume sources?
- Requires broad-band imaging and low-resolution spectroscopy.
 - Reflectance spectroscopy may sense broad bands of nitrates, hydroxides, water and organics, (e.g., species containing C-H, C=N, C=S bonds and CO₂)







Surfaces III: Asteroids

Origins of the Solar System

- Detect and map diagnostic features of many asteroid types
 - Christiansen emissivity maximum diagnoses silicate polymerization
 - Slope at shorter wavelengths gives surface porosity.
 - Comparisons to lab spectra (e.g., carbonaceous chondrites) could reveal parent reservoirs for meteorites.
- Spatially-resolved thermal emission (e.g., NEOs) can reveal variations in thermal inertia, surface properties. Photon pressures causes by emission variation influences **orbital history** of these objects.
- Requires imaging and low-resolution spectroscopy.
- Could emission from volatiles also be detected at higher spectral resolution? (e.g., volatile rich main-belt comets)



P/2013 R3 asteroid breaking up intoten pieces, taken by the Hubble Space Telescope and distributed by the European Space Agency. Photograph: D. JEWITT (UCLA)/AFP/Getty Images



Thermal model of near Earth asteroid Bennu, target of the OSIRIS-REX mission. Shape model derived from radar data. Taken from Emery et al. 2014

NASA



Surfaces IV: Distant Icy Bodies

Pluto

- Follow-up observations from New Horizons, e.g. documenting atmospheric collapse with increasing distance from the sun: CO, CH₄,
- Search for C_2H_2, C_2H_6
- Direct spectroscopy or occultation signatures.

TNOs

– Searches for atmospheres in other TNOs

Icy satellites

- Spatial variability of thermal emission
- Variable atmosphere of Triton

Rings and ring arcs of icy giants

L- through Q-band imaging and moderate- to high-resolution spectroscopy,





Comets I: Remnants of Formation

Origins of the Solar System

- Comets are the icy remnants of planet formation, compositional signatures of primitive materials reveals conditions in the protoplanetary disc
- Origin, migration and reprocessing history.
 - D/H in cometary volatiles related to terrestrial water, ¹⁵N/¹⁴N and ¹³C/¹²C.
 - Water, CO, NH₃, CH4, C_2H_2 , C_2H_6 , CH₃OH, HCN can all be measured at 3-5 μ m.

Compositional comparison of Jupiter-family comets, Della Russo et al., 2011.





Comets II: Rich L/M Spectra

Origins of the Solar System

- Comets important for volatile transport throughout forming solar system; understand source reservoir for our own planet.
- L- & M-bands:
 - Non-LTE emissions from gases, organics, some scattered sunlight.
 - Unique identification of hydrocarbons & organics.
 - Spatial mapping of source regions in cometary nucleus/coma





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Summary: Understanding the 'Near Universe'

How does our solar system work?							
Fundamental Processes	Variation of climates	Origins of the Solar System	Surface environments	Energy exchanges			
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High-resolution spectroscopy: Non-LTE emissions from atmospheres and comets; reflectance spectra for clouds; thermal emission windows on deep atmospheres and surfaces; middle-atmospheric sounding Imaging & Low-Resolution Spectroscopy: 3D sounding of planetary climate; reflectance spectra of rock/ice surfaces; thermal contrasts on satellites/asteroids; comet / asteroidal silicates.

JPIL



End



Although we're not on the road we thought we'd be on at this point, and there may be a rough ride ahead...





End



Although we're not on the road we thought we'd be on at this point, and there may be a rough ride ahead...

Let's just buckle up tightly!





End



Although we're not on the road we thought we'd be on at this point, and there may be a rough ride ahead...

Let's just buckle up tightly!

Because there are going to be great rewards at the finish

601

JPL



Supplemental Information







Giant Planet IV: External Oxygen

Fundamental Processes

How are planetary atmospheres coupled to their external environments?

- CO non-LTE emission in M-band detected on ice giants.
- External source from cometary impacts, interplanetary dust, etc.
- Spatially-resolved CO fluorescence requires high sensitivity and spatial resolution
 - Sources and sinks of giant-planet stratospheric oxygen.



Terrestrial I: Trace Species (Venus)

Variation of climates

• Drossart et al., 2007; CO₂ mapping at 4.3 µm shows a dynamic upper atmosphere





Comets III: N-band Silicate Emission

Origins of the Solar System

- Measuring the rock/ice composition of the building blocks of our solar system.
- Emissivity of Comet Tempel 1 before and after Deep Impact (Lisse et al., 2006) allows identification of silicate types
 - Diagnostic of cometary origins and thermal evolution.
- Bottom panel have silicate emissions removed, reveals carbonates, PAHs, amorphous carbon, ices.
- Requires N-band imaging and low-resolution spectroscopy.





Comet 73P/Schwassman-Wachmann 3 imaged by Spitzer in the Q-band

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